



Ionospheric sounding at the RMI Geophysical Centre in Dourbes: digital ionosonde performance and ionospheric monitoring service applications

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Introduction

- Vertical Incidence Sounding (ionosonde) measurements
- Ionosonde measurements at Dourbes
- Digisonde performance evaluation
- Applications (LIEDR)
- Summary and Outlook









Ionosphere





The **ionosphere** is the inner part of the upper atmosphere, extending from about 50 km to 1000-2000 km altitude, which is being **ionised by the solar radiation**. The major part of the ionisation is produced by solar X-ray and ultraviolet radiation and by corpuscular radiation from the Sun.



Ionosphere





In the ionosphere, free electrons occur in sufficient density to have an appreciable influence on the propagation of radio frequency electromagnetic waves. The ionospheric plasma consists of mostly H+ and He+ ions above 1000 km, O+ ions from 300 to 500 km, and molecular ions (N2+, O2+, NO+) below 200 km. Total ion densities (= electron density) range from 10^8 to 10^13 m^-3.





The vertical incidence sounding remains one of the most accurate and important ionospheremonitoring techniques. In this technique, low- and high-frequency radio waves are transmitted upward and reflected in the ionosphere at the height where the refractive index becomes zero for vertical incidence, or $sin(\phi_0)$, where ϕ_0 is the incidence angle.

The standard piece of equipment employed for the purpose is called **ionospheric sounder** (**ionosonde**), in which a transmitter and a receiver are swept synchronously in frequency, and the propagation time of the reflected signal is recorded for each of the transmitted frequencies.



Ionosonde technique:

To measure the height of signal reflection from the ionosphere at each operating frequency.

Measuring echo delay:



Ionosonde operation concept









The **ionogram** is an instantaneous record of the ionospheric conditions (above the sounder) indicated by the relationship between the frequency of the radio pulse emitted upwards and the virtual heights of echoes reflected from the ionosphere.

A typical ionogram with the key ionospheric characteristics (Wakai et al., 1987)





1957 - 1970 Ionosonde Panoramique

1970 - 1984 **Digisonde -128**



















1984 - 2011 Digisonde-256



foF2 foF1

foF1p

foE

foEp

foEs

fmin

MUF

h^{*}E

h'E

h' F2

h'Es

zmF2

zmF1

zmE

yF2

yF1

ÿΕ

. ВØ

В1

C-level

M D

f×I

6.88

4.77

4.83

3.81

3.46

7.60

4.45

1.00

21.99 3.199

3000

215

295

80

95

232 152

85

71

76

110.9

1.35

21

5

foF2	MHz	F2 layer critical frequency
foF1	MHz	F1 layer critical frequency
foF1p	MHz	Predicted value of foF1
foE	MHz	E layer critical frequency
foEp	MHz	Predicted value of foE
fxl	MHz	Maximum frequency of F-trace
foEs	MHz	Es layer critical frequency
fmin	MHz	Minimum frequency
MUF(D)	MHz	Maximum usable frequency
M(D)	-	M(D) = MUF(D)/foF2
D	km	Distance for MUF calculation
h'F	km	Minimum virtual height of F trace
h'F2	km	Minimum virtual height of F2 trace
h'E	km	Minimum virtual height of E trace
h'Es	km	Minimum virtual height of Es trace
zmF2	km	Peak height of F2-layer
zmF1	km	Peak height of F1-layer
zmE	km	Peak height of E-layer
yF2	km	Half thickness of the F2 layer,
yF1	km	Half thickness of the F1 layer
уE	km	Half thickness of E layer
B0	km	IRI thickness parameter
B1	-	IRI profile shape parameter
C-level	-	Confidence level: 1 (highest)



100 200 400 600 800 1000 1500 3000 [km] D MUF 7.6 7.7 8.0 8.6 9.3 10.4 13.7 22.0 [MHz] DB049_2003154110005+MMM / 220fx128h 50 kHz 5+0 km 2x3 / DGS-256 (049-049) 50+1 N 4+6 E





April 2011 -

- Type: Lowell Digisonde-4D Location: Dourbes (50.1°N, 4.6°E) URSI code: DB049
- Cadence: 5 min















2011 -Digisonde-4D















- general increase in the critical frequencies with solar activity
- foF1 and foE in phase w/ the solar zenith angle
- foF2 in antiphase w/ the solar zenith angle (the F2 winter anomaly)
- F1 may disappear in winter
- mid lats: foF2 diurnal max at midday in winter, late afternoon in summer
- low lats: foF2 diurnal max may also occur in the evening













(**A**/**B**): Successfully autoscaled ionograms







(A/B): Autoscaling failures –
(A) partial, E-layer
parameters only scaled, and
(B) completely unscaled, an
example of severely
depleted ionosphere during
a geomagnetic storm.

(**C**/**D**): Gap occurrence (due to interference) - smaller gaps successfully ignored/interpolated, larger gaps falsely inter-/extrapolated resulting in the automatic layer trace being truncated prematurely (D).

(**E**/**F**): Incorrect autoscaling of the h'F2 virtual height.































(A/B): Histograms of the foF2 and h'F2 errors. (C): Relative cumulative foF2 and h'F2 error distributions.

Error bounds (95% probability): foF2 (-0.75,+0.85), foF1(-0.25,+0.35), foE(-0.35,+0.40), h'F2(-68,+67), h'F(-38,+32), h'E(-26,+2), M3000F2(-0.55,+0.45)





• Automatic scaling availability

available in 94-98% cases for all characteristics except foF1 (89%)

Autoscaling accuracy

for some characteristics (most notably for foF2 and M3000F2) the magnitude of the residual error (autoscaled minus manually-scaled values) **varies in local time, season and solar activity**

Influence of geomagnetic activity/storms

Although geomagnetic storms seem to affect the autoscaling, the overall results about the influence of geomagnetic activity remain **inconclusive**

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Overall, the automated ionogram processing/scaling has demonstrated **sufficiently good performance** that allows the utilisation of the instantaneous ionospheric sounding data for operation of a monitoring system





Physical & Mathematical background

Roval

Institute

 $\mathfrak{I}_{i}(h) = \mathfrak{I}_{i}(H_{i}, N_{mi}, h_{m}F_{2}; h)$ ion density profile (topside) The height profile of the electron density calculated via the following 'reconstruction' formula: $\mathfrak{I}_{e}(h) = \mathfrak{I}_{O+}(H_{O+}, N_{mO+}, h_{m}F_{2}; h) + \mathfrak{I}_{H+}(H_{H+}, N_{mH+}, h_{m}F_{2}; h) \leftarrow \text{plasma density profile (topside)}$ The unknowns $(H_{0+}, H_{H+}, N_{m0+}, N_{mH+})$ determined from the following system of equations: $N_{mO_{+}} + N_{mH_{+}} = N_{m}F_{2}$ 🗲 plasma quasi-neutrality (peak) $H_{H_{+}} = \left(\mu_{O_{+}} / \mu_{H_{+}} \right) \xi H_{O_{+}}$ O⁺/H⁺ scale heights relation $\Phi_{t} = \aleph_{O_{+}}(H_{O_{+}}, N_{mO_{+}}, h_{m}F_{2}) + \aleph_{H_{+}}(H_{H_{+}}, N_{mH_{+}}, h_{m}F_{2})$ integrated topside densities $\mathfrak{I}_{O_{+}}(H_{O_{+}}, N_{mO_{+}}, h_{m}F_{2}; h_{tr}) = \mathfrak{I}_{H_{+}}(H_{H_{+}}, N_{mH_{+}}, h_{m}F_{2}; h_{tr})$ O+/H+ ion transition level Atitude H_{O+}, H_{H+} - the O⁺ and H⁺ ion scale heights N_{mO+} , N_{mH+} - the O⁺ and H⁺ ion maximum densities UTL μ_{O+} - the O⁺ ion mass (atomic mass = 15.9994 amu, 2.6567625437×10⁻²⁶ kg) $\mu_{H\scriptscriptstyle +}$ - the H+ ion mass (atomic mass = 1.00794 amu, 1.6737235385×10^{-27} kg) slab thickness Hsc ξ - the vertical 'scale height' corrector, $\xi = \sin[\arctan(2\tan\varphi)], \varphi$ -geom. latitude hmF2 h_{tr} - the upper, O⁺/H⁺ ion transition level Φ_{t} - the measured topside TEC (above hmF2) NmF2 $\aleph_{mO_{+}}$, $\aleph_{mH_{+}}$ - the integrated topside O⁺ and H⁺ ion densities Number density



Local ionospheric plasma density specification in real time















http://swans.meteo.be/ionosphere/liedr

Local ionospheric plasma density specification in real time Meteorological

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Modern GNSS-based applications demand high precision – simultaneous real-time observation of several characteristics essential (incl. solar/geomagn. activity and 'derivative' measures e.g. ionospheric slab thickness)

Electron density reconstruction technique - reliable, easy to maintain and upgrade. It is important that new measurements can be obtained and processed rapidly, which in turn provides higher resolution in the results. Possibilities for extension to regional ionosphere monitoring (for regions with dense ionosonde networks, e.g. Europe; alternatively, using empirical/model foF2 maps)

Research applications: further understanding the ionospheric morphology, validating existing ionospheric models. Suitable for investigating local ionospheric storm-time development. However, for better identifying and observing a storm, it is necessary to include geomagnetic measurements

Operational applications: ionospheric/space weather monitoring, research & modelling -- to improving comm/nav systems performance (incl. HF propagation and ray tracing, adverse ionospheric effects warnings/mitigation)

Further developments – reconstruction using variable scale height profilers, improving the ion transition height model